

Issues to be considered for strategic adaptation to climate evolution Is atmospheric evaporative demand changing?

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Abstract

The predicted developments in climate are region-specific and adaptation can only be successful considering the regional characteristics with its diverse technical, environmental, economic and social implications. Beyond some obvious adaptation strategies in response to emerging environmental constraints for example there are many more “basic” challenges below “the surface”. One of the key concerns for many regions is the availability of water and how increasing temperature will drive the evaporative demand of the atmosphere. For this, individual regions need to be analysed to quantify possible associated risks. This paper will address differences in regional water relations of grape growing areas in different parts of the world as a basis to address the points listed above.

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Introduction

Climate change effects on the terrestrial water cycle show regional differentiated patterns. While temperature is increasing in many world grape growing regions (Jones *et al.*, 2005; Schultz and Jones, 2010; Webb *et al.*, 2012; Hannah *et al.*, 2013; Tóth and Végvári, 2016) precipitation patterns can vastly differ between regions and can show substantial temporal variations (between and within years) (IPCC, 2014). From rising temperatures it is mostly assumed that water holding capacity of the atmosphere will increase in the future as a function of the Clausius-Clapeyron law (Krysanova *et al.*, 2008) which predicts an increase in the saturation vapour pressure of the atmosphere of 6-7 % per degree Celsius. As a consequence, a simultaneous increase in potential evapotranspiration (the amount of water that could potentially be evaporated from soils and transpired by plants due to changes in climatic factors such as temperature, vapour pressure deficit, radiation and wind speed, ion of water from the soil and transpiration of water from plants, ET_p) is assumed in many cases, which would alter soil and plant water relations. However, the same underlying principles also predict an increase in precipitation by 1-2 % per degree warming (Farquhar and Roderick, 2007). Additionally, model predictions for many regions forecast altered precipitation patterns and thus in combination with the possibility of increased ET_p, farmers around the world fear an increase in the likelihood of water deficit and the availability of water for irrigation.

However, the large spatial and temporal variability in precipitation patterns between regions preclude generalizations in predicted consequences with respect to soil and plant water status development. Especially the temporal variability may mask longer-term trends in the development of ET_p and consequently soil and plant water status (vVan Leeuwen *et al.*, 2010). Additionally, the focus on the developments within a growing season (spring-summer) in many studies may miss decisive effects occurring during the “off-season” (winter-early spring) but having substantial carry-over effects into the season.

Evaporation is driven by changes in temperature, humidity, solar radiation and wind speed and contrary to expectations due to climatic changes, there have been reports on a reduction in evaporative demand worldwide (Farquhar and Roderick, 2007). In many cases this has been related to a decrease in solar radiation observed for many areas on earth including wine growing regions in Europe until the beginning

of the 80th (global dimming, (Wild *et al.*, 2005; Hofmann and Schultz, 2010)) of the last century. However, ET_p in some areas has continuously increased which suggests that changes in the aerodynamic component must have more than offset the decrease in radiation over that part of the observed time span (Schultz and Hofmann, 2016). For some regions in Germany, wind speed and vapour pressure deficit (VPD) of the atmosphere have increased in the past and contributed to changes in evapotranspiration (Bormann, 2011) but this is not in agreement with a worldwide observed decrease in wind speed and pan evaporation (Farquhar and Roderick, 2007; McVicar *et al.*, 2012).

These conflicting observations depending on climate classification, country or region, make it necessary to analyze grape growing regions with respect to developments in ET_p and precipitation patterns much more in detail in order to make predictions with respect to an increased risk in terms of water shortage. There is a general lack of studies analyzing the past development in ET_p and precipitation for different wine growing regions across the planet in order to answer the question whether the threat for sustained drought will increase. When ET_p was set to increase in a future climate scenario, substantial reductions in pre-dawn leaf water potential resulted when a dynamic physiological grapevine water model was used (Lebon *et al.*, 2003) to estimate water consumption (Schultz and Lebon, 2005). However, the large spatial and temporal variability in precipitation patterns between regions preclude generalizations in predicted consequences with respect to soil and plant water status development.

Water limited worlds versus energy limited worlds

Those parts of the earth where evaporative demand exceeds supply (rainfall), like many Mediterranean-type climatic regions, are very different from those parts of the world where rainfall exceeds evaporative demand, like for example Germany or many French grape growing areas. In the latter areas there is drainage to aquifers and runoff and to rivers, and evaporation rate largely depends on the available energy and especially the radiation received. In water-limited regions, there is an excess of energy (e.g. solar radiation), and the actual evaporation rate can be close to the rainfall (Farquhar and Roderick, 2007). Grape growers from these different parts of the world have a very different view on their environment. The distinction between water limited versus energy limited worlds is not completely consistent because winters for example in water

limited areas will, in a lot of many cases, be part of the energy limited “world” in Fig. 1 based on Budyko (1974) and a conceptual analysis of Farquhar and Roderick (2007). Following this analysis, the actual evaporation rate, E_a , must be less than or equal to evaporative demand, ET_p , and also less than or equal to precipitation, P (Fig. 1). The water-limited regions or the water limited part of the season (which could be part of both general areas) are on the left, and the energy-limited regions (or parts of the season) are on the right of the figure.

Material and Methods

The base of possible changes in ET_p is a change in temperature which has been observed in many regions. One of the largest collections of climate data (temperature, rainfall, sunshine hours but no data on ET_p) from grape growing regions has been published in Gladstones (1992) based on observations for different time periods depending on the source and the availability of data. Four of these data sets from different regions in the northern and southern hemisphere with a strong reputation for growing Cabernet Sauvignon (Bordeaux, Napa Valley, Coonawarra and the Barossa Valley) were extracted from Gladstones (1992) and compared to data from the same stations (or at least very close) for the last 25 years since 1990 (end of observational period in Gladstones book) to get a feeling of investigate the magnitude of changes in temperature which has

occurred over that time span. Figure 2 shows the differences in average monthly temperature for the growing season months.

It is obvious from Figure 2 that in all regions shown, temperature has increased albeit to a different extent, which in principle would satisfy the Clausius-Clapeyron relationship which predicts an increase in the saturation vapour pressure of the atmosphere of 6-7 % per degree Celsius, thus increase ET_p if air moisture would not change.

In order to evaluate different grape growing regions with respect to observed changes on precipitation patterns and ET_p and in order to validate or disprove general observations on changes across the planet (Farquhar and Roderick, 2007), the data of five wine-growing areas in four countries in the Northern and Southern hemisphere across a large climatic transect were analyzed. Climatic data for this analysis were provided by the German Weather Service (Deutscher Wetterdienst) for the location Geisenheim in Germany (50,0° N, 8° E) in a temperate climate, the French INRA CLIMATIK, Agroclim project for the locations Dijon, Burgundy (47,2° N, 5,2° E), temperate climate, and Avignon (43,9° N, 4,9° E) in a Mediterranean climate, the US California data provision system on integrated pest management for Oakville, Napa Valley, CA (38,3° N, 122,3° W), a Mediterranean climate situation, and the Australian Government, Australian Bureau of Meteorology, for

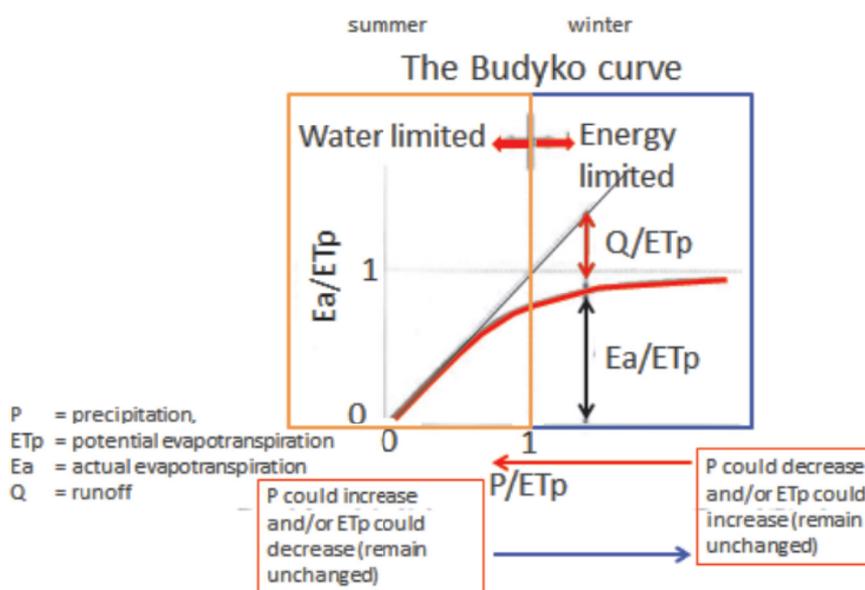


Figure 1 - Inter-relationship between average precipitation (P), actual (E_a) and potential (ET_p) evapo-transpiration and runoff (Q) and how season and climate change could affect this inter-relationship depending on the region.

Grape growing areas are represented in both water and energy limited areas and the effect of climate change might be substantially different for different parts of the world. The original curve is known as the Budyko curve (Budyko 1974) and the presented figure is an adaptation from Farquhar and Roderick (2007) in an extended version.

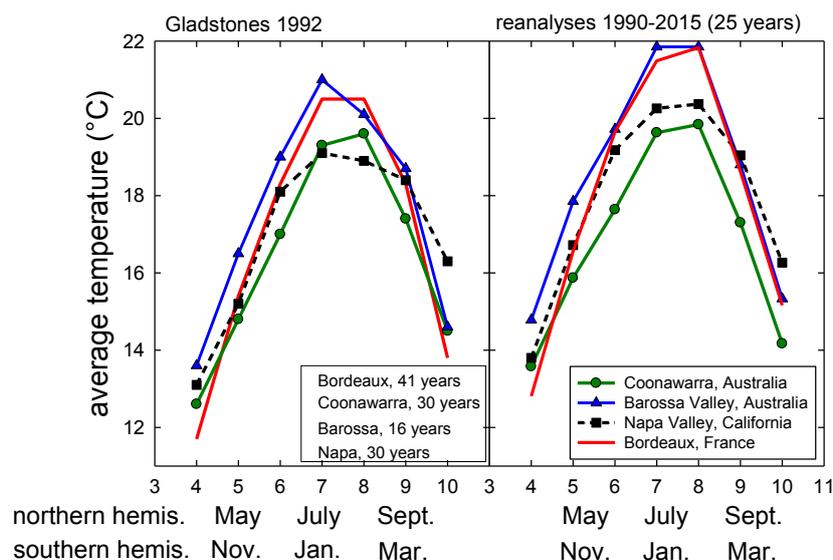


Figure 2 - Average monthly temperatures for four different grape-growing regions with a reputation to grow Cabernet- Sauvignon.

Left panel; data extracted from Gladstones book (1992), right panel; data from the same stations where possible or for stations close by for the last 25 years (1991-2015). Data sources were Météo Bordeaux, US California data provision system and the Australian Bureau of Meteorology.

Williamstown, Adelaide hills (34,7° S, 138,9° E). Data were seasonally separated into precipitation and ETp “summer” for the growing season (May-October for the northern hemisphere, October-May for the southern hemisphere), which in agro-meteorological terms is defined as the “hydrological summer” (Bormann, 2011), and the “off-season” (November-April for the northern hemisphere, April-November for the southern hemisphere), the “hydrological winter”. In the case of the German data, predictions for precipitation rates and ETp were used based on model-outputs of a regionalized version of the STARII model of the Potsdam Institute of Climate Impact (Orlowski *et al.*, 2008). STARII constructs time series from 2007-2060 by resampling of observed weather data according to trend informations of the Global climate model ECHAM5/OM (A1B) (Jacob, 2005). This approach provides physical consistency of the combination of the weather variables and is in close agreement compared to the statistics of observed climatology (Orlowski *et al.*, 2008).

Results and discussion

The general expectation, which is also very prevalent in the popular press, that as the world warms because of increased greenhouse forcing there will be a widespread increase in evaporative demand has been challenged by data proving the contrary and by a lack of scientific basis put forward by several scientists

(see discussion by Farquhar and Roderick, 2007). Peterson *et al.* (1995) were the first to publish the results from 190 sites in the former Soviet Union, where they found decreasing pan evaporation rates in the European sector, a decline in Siberia, and no trend in the Asian part. Since then many other reports from different parts of the world have been published but none has explicitly looked at grape-growing regions.

Observed and predicted summer trends for areas in Europe and California

Figure 3A shows observed (calculated according to Penman-Monteith) and predicted changes in ETp during the growing season (May-October) for the temperate wine-growing region of the Rheingau (Geisenheim, Germany, 50.0° North, 8° East) from 1958 until 2060 (Schultz and Hofmann, 2016). To smooth out temporal variability, 10-year running mean values were used. There is a clear increase in the difference between ETp and precipitation rate during the growing season already observed during the past 55 years and this development will continue in the future as predicted using a regionalized version of the STARII model (Orlowsky *et al.*, 2008) (Fig. 3A). A similar increase in ETp was also observed for the Mediterranean region near Avignon, France, since the mid-seventies of the last century, but with no observed change for about the last 20 years (Fig. 3B). Available data for the Napa Valley in California show that ETp has not changed for approximately 30 years

despite concomitant observations on rising temperatures.

Obvious from Figure 3A are the cyclic patterns of both ETp and precipitation rates, both for the period of observation and the projections until 2060. These cycles may be related to solar cycles which have been made partly responsible for the warming during the first half of the last century but not during the second half (Stott *et al.*, 2003). However, there is some uncertainty on whether these cycles do continue to have an impact on the temporal development of warming on earth and consequently on evaporation (Stott *et al.*, 2003) but the data do show that variability and the development of extremes will become more likely despite cyclic variations (Fig. 3A) (IPCC, 2014). These cycles have an important effect on how climate change is perceived by humans since they can somewhat mask long-term trends (when precipitation is increasing or ETp is decreasing for several years) or on the contrary suggest a speed-up in these trends (Fig. 3A).

Precipitation trends in Avignon have undergone some fluctuations but there was no distinct decrease observed, similar to summer precipitation in the Napa Valley, albeit on a much lower level (Fig. 3B). If ETp predictions for the cool climate area of Germany (50°

North) would be correct, then summer ETp values by the middle of the current century would be similar to Avignon (43,9° North) in the seventies at lower precipitation rates.

Observed trends for Australian and California regions (summer and winter)

Analyzing data from one Australian region, Williamstown in the Adelaide Hills, it is obvious that neither ETp nor precipitation have changed substantially over the time period of available data confirming other data from Australian sites (Roderick and Farquhar, 2004) (Fig. 4). The long-term data set from Williamstown shows that ETp decreased between the seventies and the nineties during both winter and summer before increasing again to the early ETp values. This might have been related to the phenomenon of global dimming, a reduction in solar radiation observed in many areas during that particular period caused by increased cloudiness and aerosols (Wild *et al.*, 2005; Hofmann and Schultz, 2010). Precipitation rates also show no clear trend with a slight decrease during winter for the Adelaide Hills (left panel, Fig. 4). Similarly, ETp during winter and summer of the Napa Valley location did not change appreciably (Fig. 4), yet winter precipitation has almost been halved over the

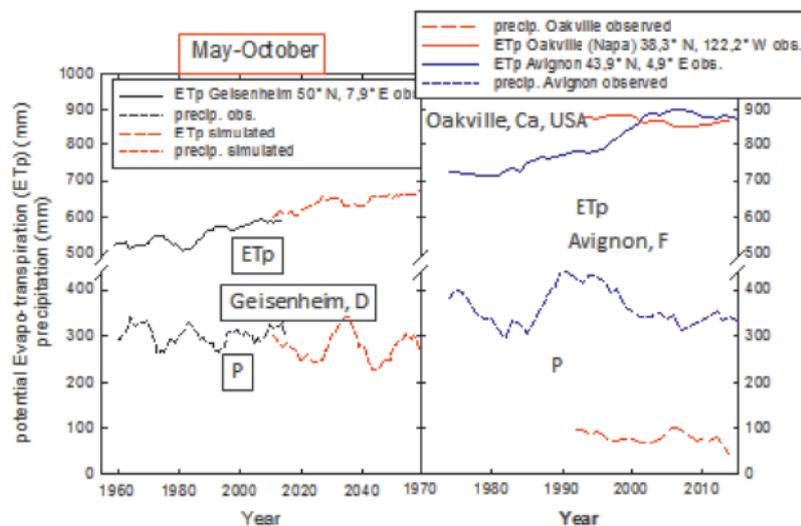


Figure 3 - Observed and simulated precipitation and potential Evapotranspiration for the hydrological summer (May-October) for Geisenheim in the Rheingau region (Germany, 50° North; 8° East) (left panel).

Potential Evapotranspiration rates for the observed time period (1958-2013) were calculated according to Penman-Monteith. Simulations were conducted with the STARII model of the Potsdam Institute of Climate Impact using the medium realization run (Orlowski *et al.* 2008) (adapted from Schultz and Hofmann 2016). In the right panel, observed ETp and precipitation data are shown for two Mediterranean type climate locations, one in Avignon, France, the other at Oakville in the Napa Valley, California. Data show 10-year running mean values. Observed data were from the Deutsche Wetterdienst, Germany, the French INRA CLIMATIK, Agroclim database and the US California data provision system on integrated pest management at the University of California, Davis.

past 25 years, moving the area from an energy limited towards a water limited part on the Budyko curve (Fig. 1). Despite of a “natural” focus on the developments within the growing season, changes in the water budget during the “off-season” seem to become more important (Fig. 4 left panel). Regardless of the fact that during winter and spring precipitation rates are exceeding ETp, the “gap” between these two factors determining the soil water balance is decreasing in some areas (IPCC, 2014). This suggests that for this particular region winter precipitation will eventually be matched by winter ETp with important consequences for the amount of water stored in the soils at the beginning of the growing season. It may also have consequences for the use of cover crops during the winter.

The phenomenon that ETp remains stable or decreases in many regions even in the post-global dimming period has been related to different combinations of effects, yet the most pronounced effect seems that the wind speed in many areas has decreased (Farquhar and Roderick, 2007). A recent paper on the situation in China showed that wind speed has declined by 25-30 % since the nineties (Liu *et al.*, 2014) and a decrease of similar magnitude has been observed for the Cape region in South Africa (Hoffmann *et al.*, 2011) and are implicated in the worldwide decrease in evaporative demand (McVicar *et al.*, 2012). Data on wind speed are not easily

available, but over the same time period, wind speed has not changed in several German regions (data not shown) and in some even an increase has been observed (Bormann, 2011), which could be part of the explanation of different trends for different areas.

Observed trends for cool climate regions in Germany and France (winter and summer)

Aside of Mediterranean-type, low summer rainfall climates (water limited) with a more or less continuous decline in water availability over most of the growing season, temporary water deficits also commonly occur in temperate, summer rainfall regions, specifically on vineyard sites with shallow soils and low water holding capacity (i.e. vVan Leeuwen *et al.*, 2010). As compared to an irrigated vineyard situation in moderate or even hot climates, the natural cycles of stress and relieve can be much more pronounced albeit completely unpredictable in frequency, duration and severity in these areas and are naturally part of the ‘terroir’ and the year to year variation in wine quality. Most classic European grape growing regions are unirrigated and examples are given for two classical cool climate regions and the observed trends in ETp and precipitation during winter and summer (Fig. 5). Despite being classified as cool climate regions, both precipitation and ETp differ vastly. Geisenheim has higher ETp than Dijon in winter (Fig. 5A) and up to the nineties this was

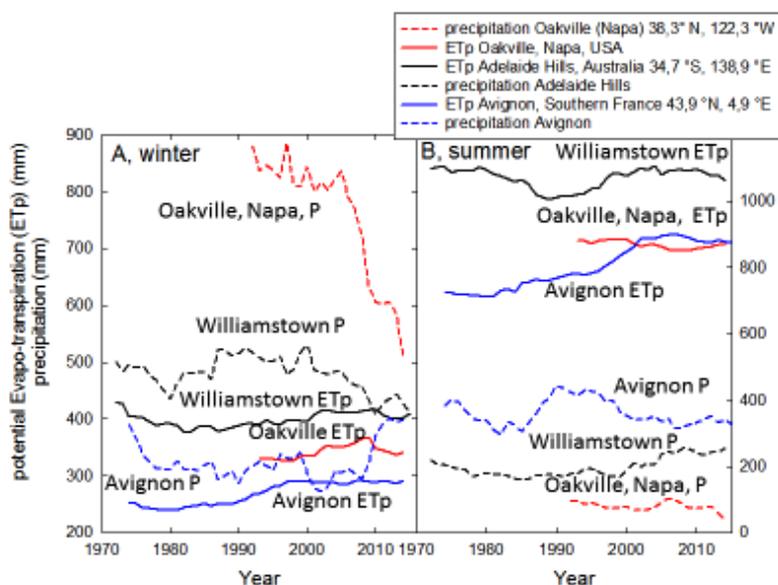


Figure 4 - Observed precipitation and potential Evapotranspiration for the winter (left panel) and summer periods (right panel) for Oakville, Napa Valley, California (USA, 38,3° North, 122,3° West) and Williamstown in the Adelaide Hills (Australia, 34,7° South, 138,9° East).

Avignon data from France have been added to illustrate how different regional situations are in broadly defined “Mediterranean climates”. Data show 10-year running mean values. Observed data were from the US California data provision system on integrated pest management at the University of California, Davis and the Australian Bureau of Meteorology, Australian Government and the French INRA CLIMATIK, Agroclim database.

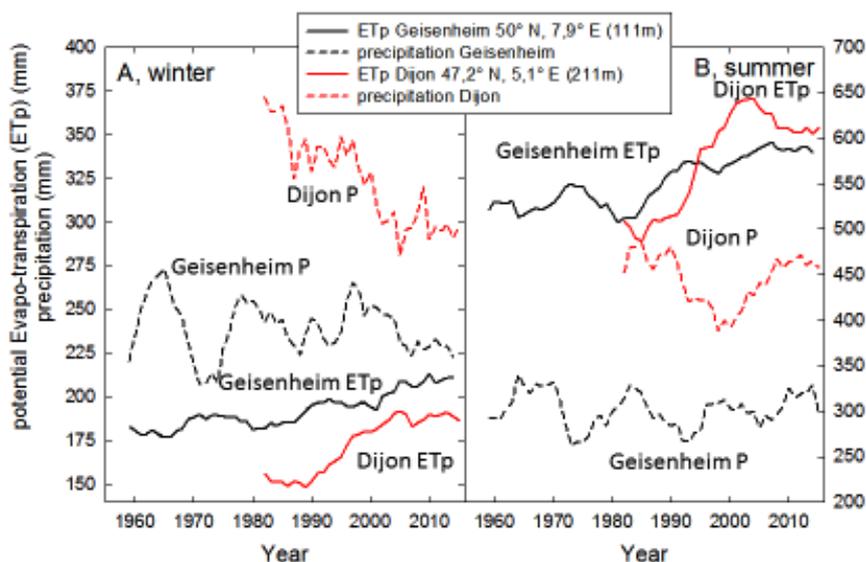


Figure 5 - Observed precipitation and potential Evapotranspiration for the winter (left panel, A) and summer periods (right panel, B) for Geisenheim, Germany (50° North, 8° East) and Dijon, Burgundy, France (47,2° North, 5,1° East). Data show 10-year running mean values.

Observed data were from the Deutscher Wetterdienst, Germany, and the French INRA CLIMATIK, Agroclim database.

also the case for summer (Fig. 5B). Geisenheim shows a continuing increase in ETp over the past 60 years in both winter and summer, whereas Dijon in Burgundy showed a strong increase starting in the nineties for both winter and summer with no change or even a decline over the past 10-15 years during the summer months (Fig. 5B). Precipitation follows a cyclic trend in all regions and in all seasons with a strong decrease in winter precipitation in Dijon over the last 35 years (Fig. 5A). In general Precipitation and ETp are inversely correlated which would be according to theory (Farquhar and Roderick, 2007).

Conclusions

The data show that generalisations with respect to global developments are not possible and that each individual region needs to be analysed with respect to observed trends and also with respect to expected developments (Hofmann *et al.*, 2014). The reasons for different developments in ETp seem to be complex and little understood. Trends might also be influenced by the drawing of moisture from water bodies which could balance the increases in temperature. According to the Budyko hypothesis, change in actual evaporation in dry regions is dominated by change in precipitation rather than potential evaporation. In humid regions, such as the cool climate examples given here, the change in actual evaporation is controlled by change in potential evaporation rather than precipitation, which would mean that the development of water deficit would become more likely in the future. Of all

regions analysed, none has shown a continued decrease in ETp or an increase in precipitation as observed for other parts of the world (Farquhar and Roderick, 2007). Rising CO₂-concentration with its effect on stomatal closure and thus potential reduction in water use may also play a role in changes in the balance between precipitation and ETp (Gedney *et al.*, 2006).

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